

*State of the Arctic: Perspectives from 4 Corners of the North*

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M Welker ripping it up on  
Prince William Sound


## Current Status of Polar Regions



**Change in the  
Polar Regions**



Winter is key to vegetation change



Herbivores modify  
biogeochemical  
responses to warming



High Arctic C cycling is more  
important than we thought

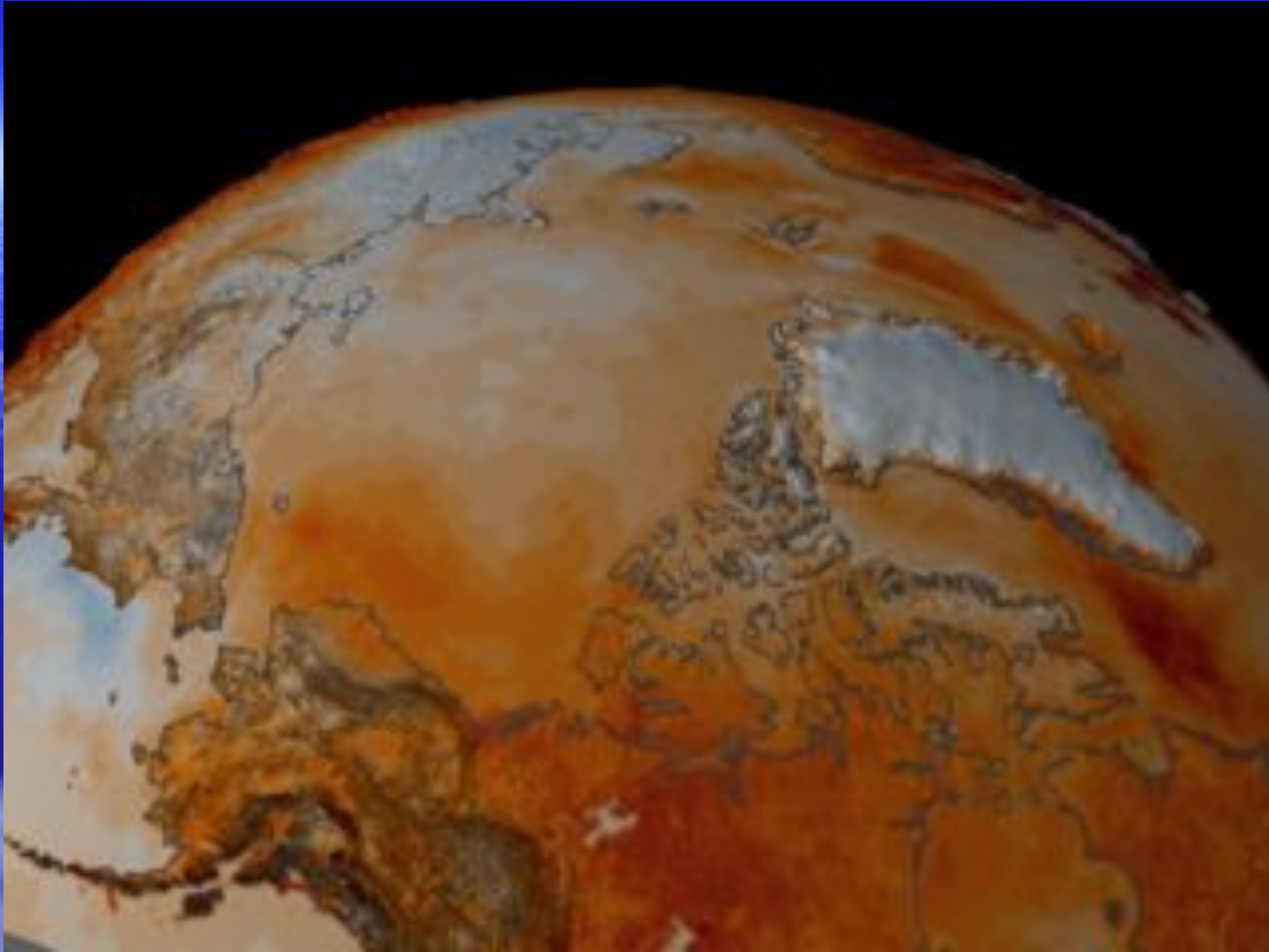


Thawing permafrost may be an important  
source of water for northern plants

# Overview

- **Background**
  - Changing climates in the north & IPY Efforts
- **Perspectives from the Four Corners**
  - **Winter**
    - Year-long considerations and fall-winter transitions
    - Controls on vegetation change
    - Snow arrival and spring vegetation growth
  - **Herbivores in the Arctic-controllers and responders**
    - Climate warming interactions
  - **High Arctic C cycling**
    - Temperature and water controls on CO<sub>2</sub> exchange
    - Soil C pools and recalibrations of Arctic C pools
    - Ancient C-modern C cycle interactions
  - **Ecohydrology and change**
    - Water sources in boreal and tundra systems
    - Feedbacks to atmospheric water vapor-the forgotten greenhouse
- **Summary**

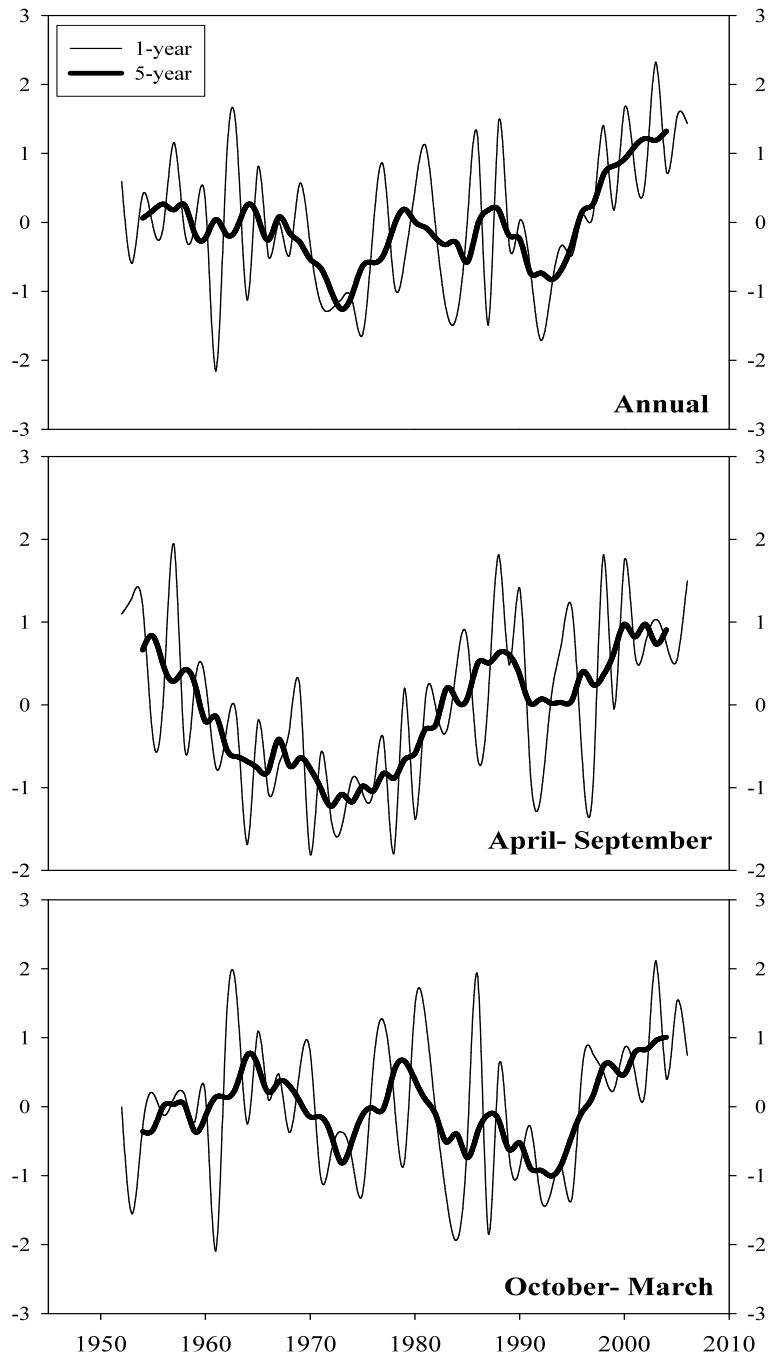




The Arctic is changing and is warming in many locations



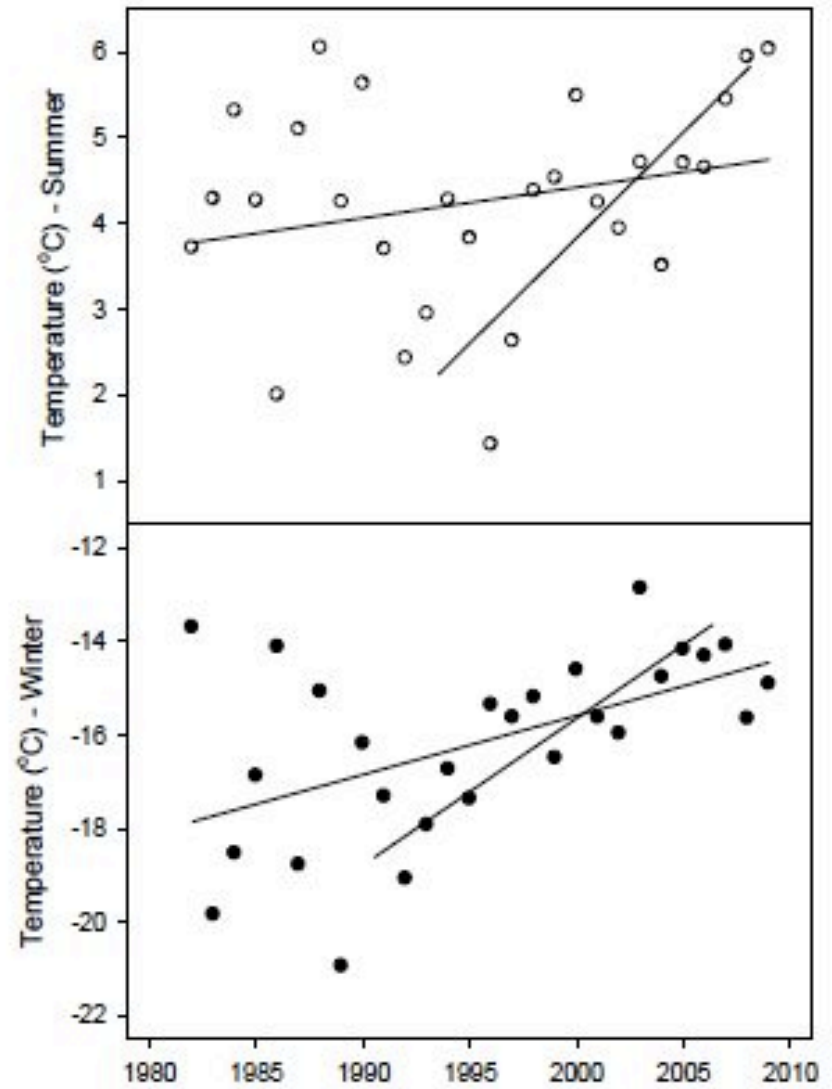
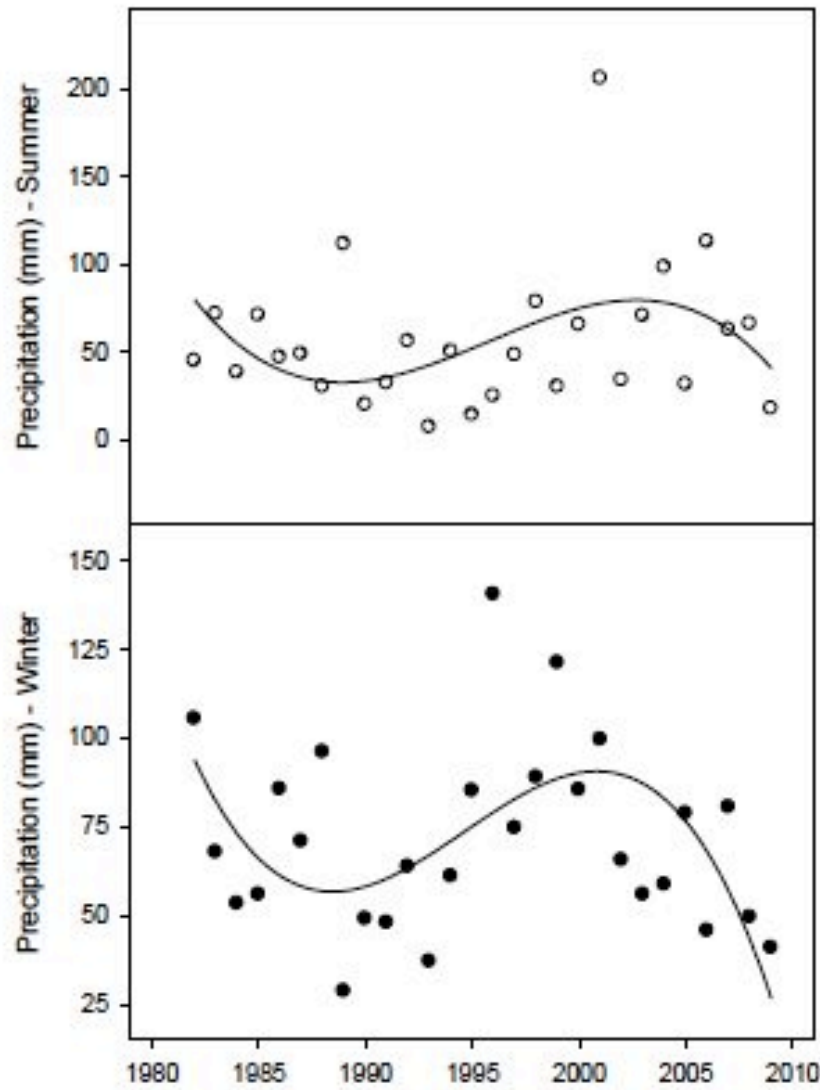
Temperature-°C anomalies



**Warming in NW  
Greenland has been  
pronounced since the  
early 1990's preceded  
by cool periods in the  
1970's and in the late  
1980's**



# Recent Precipitation and Temperatures in NW Greenland



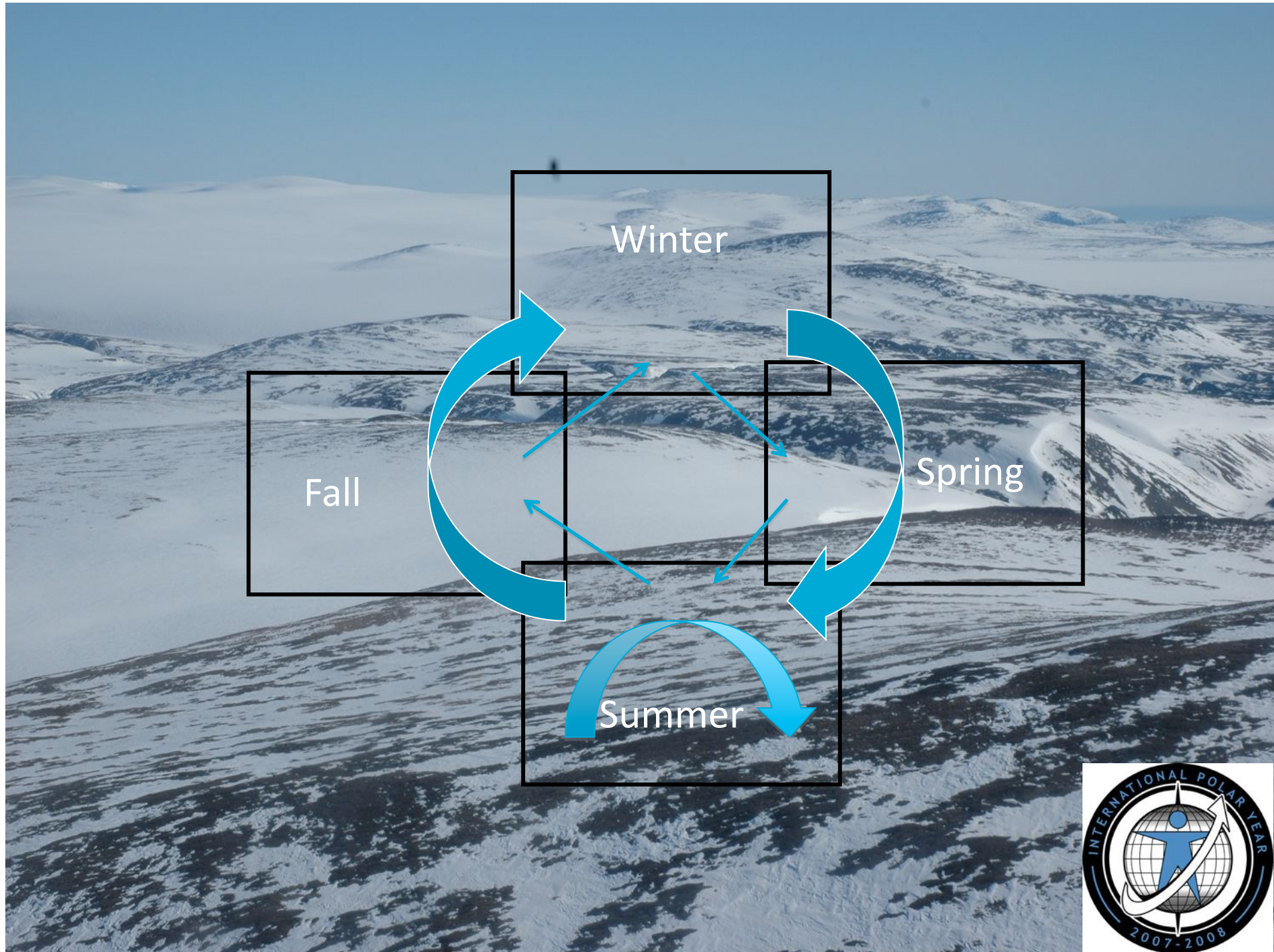


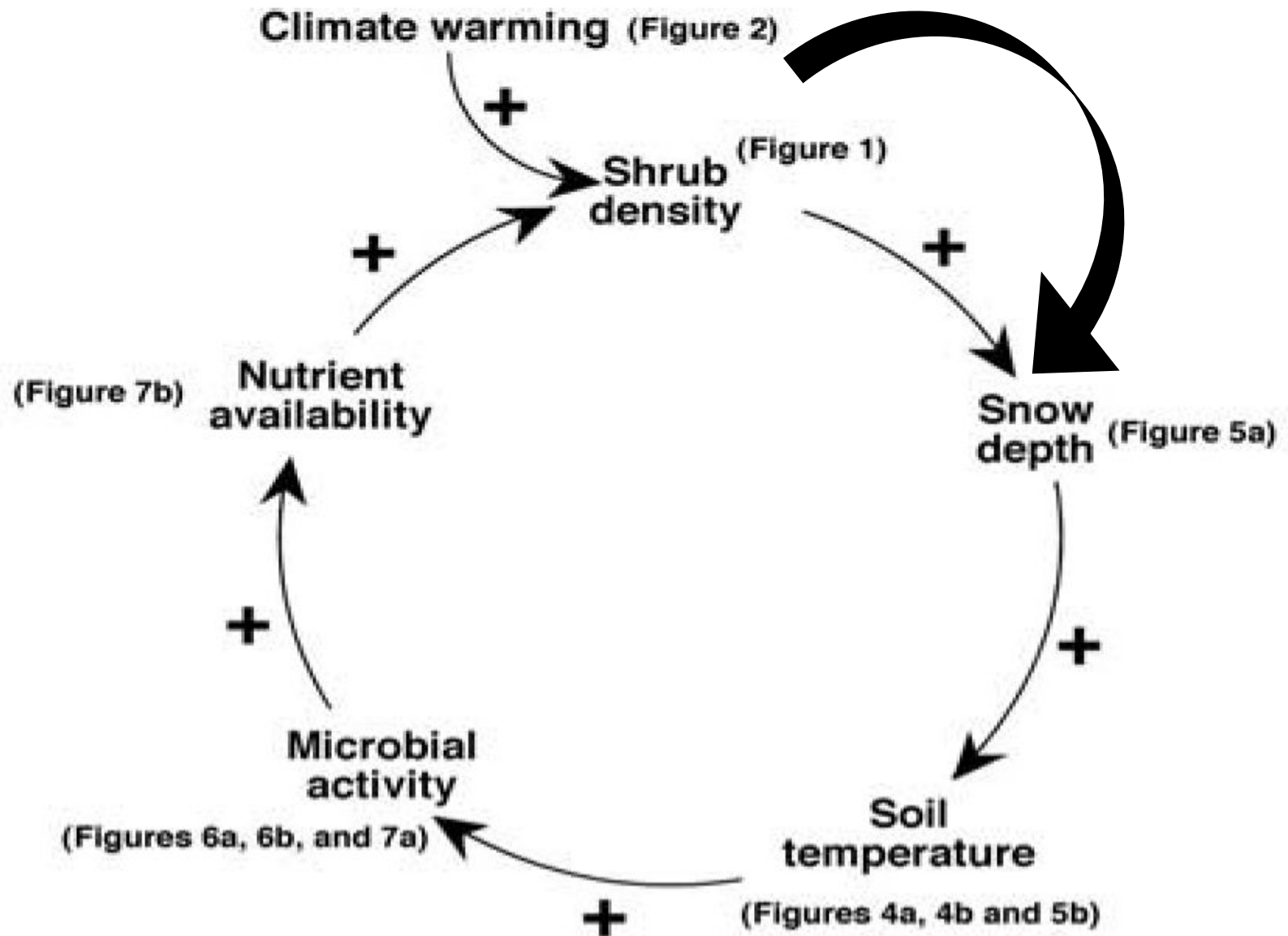
Thule, NW Greenland



Toolik Lake, Alaska

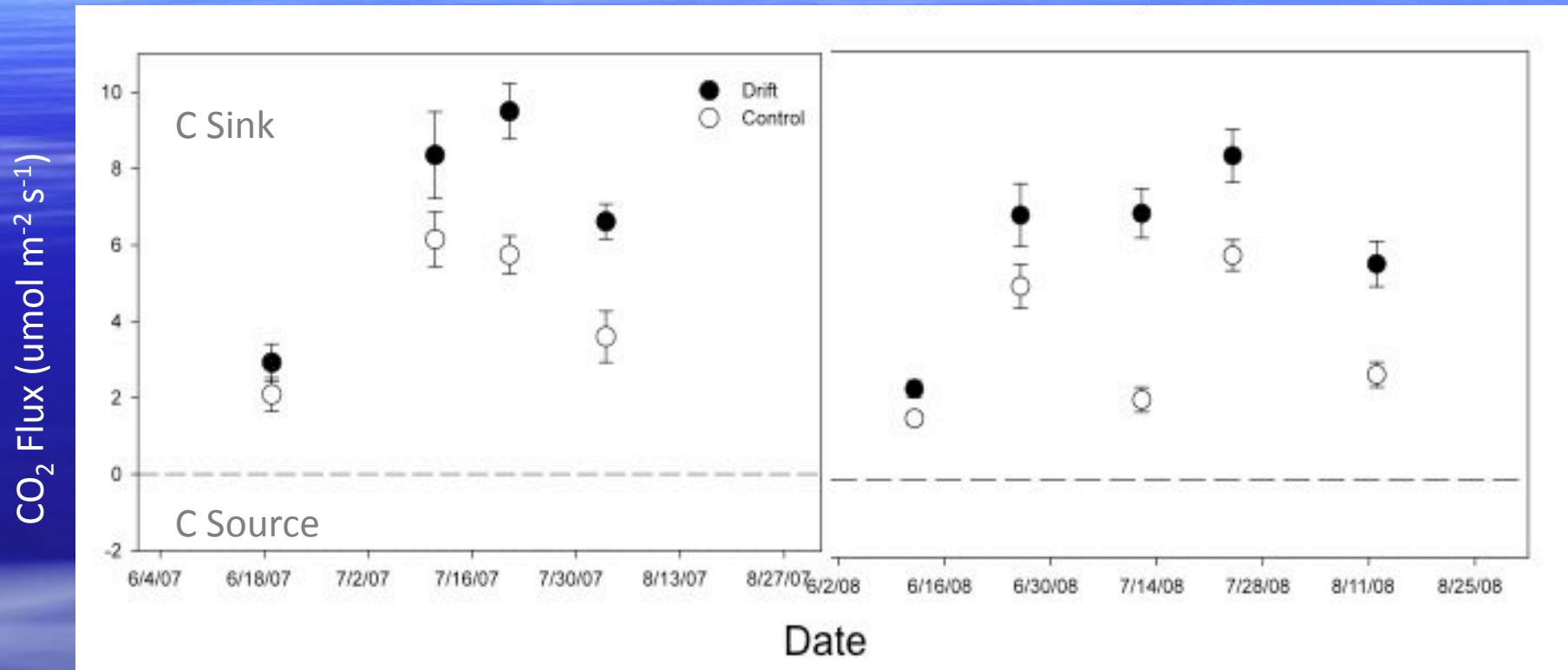






Fahnestock et al. 1998, 1999, Schimel et al. 2004, Sturm et al. 2005, Welker et al. 2005

# Net Ecosystem Exchange



Intermediate increases in snow leads to stronger net CO<sub>2</sub> sink activity-greater C sequestration

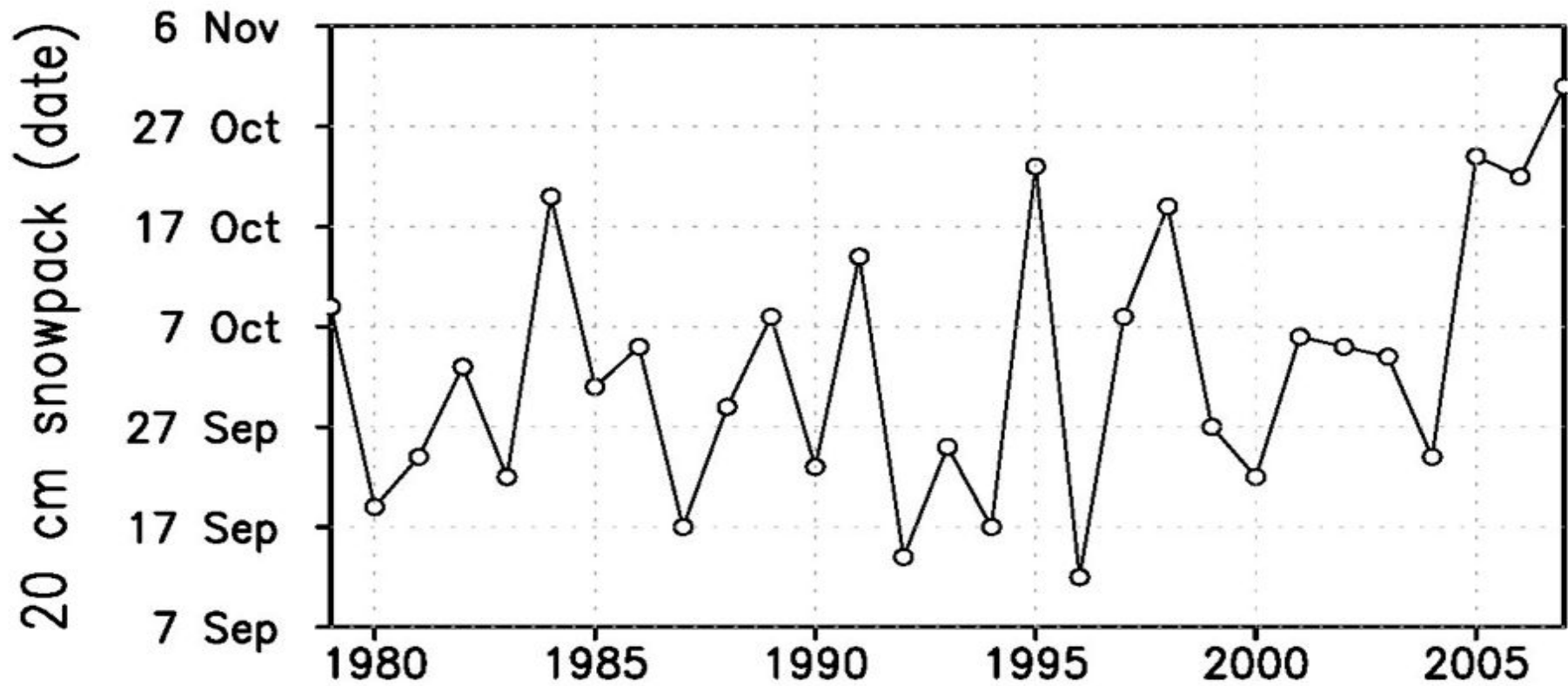
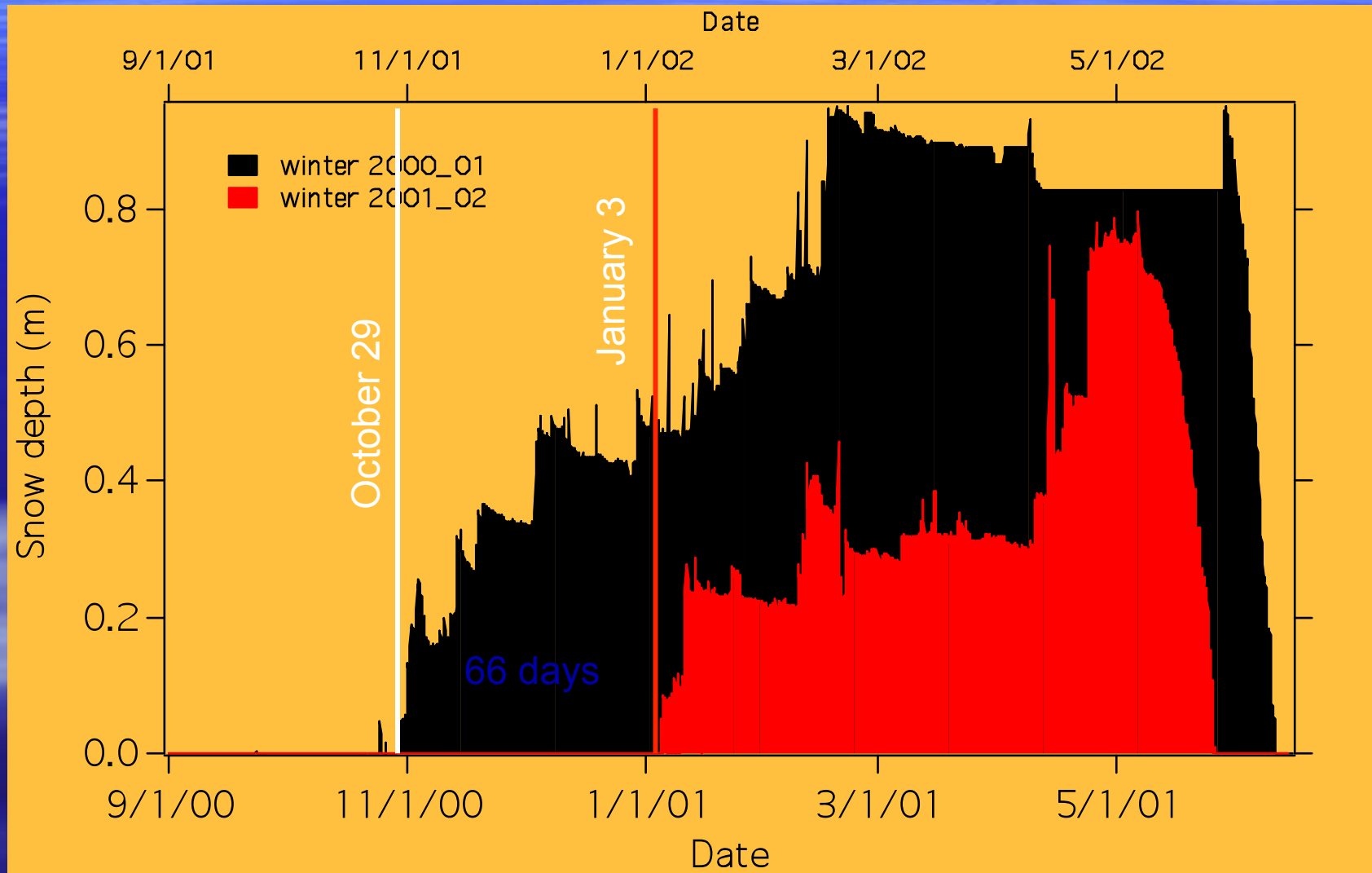


Figure 2. The arrival of 20 cm deep snow at Toolik Lake. Calculated using SnowModel (Liston and Elder 2006a,b) for 1979-2007.

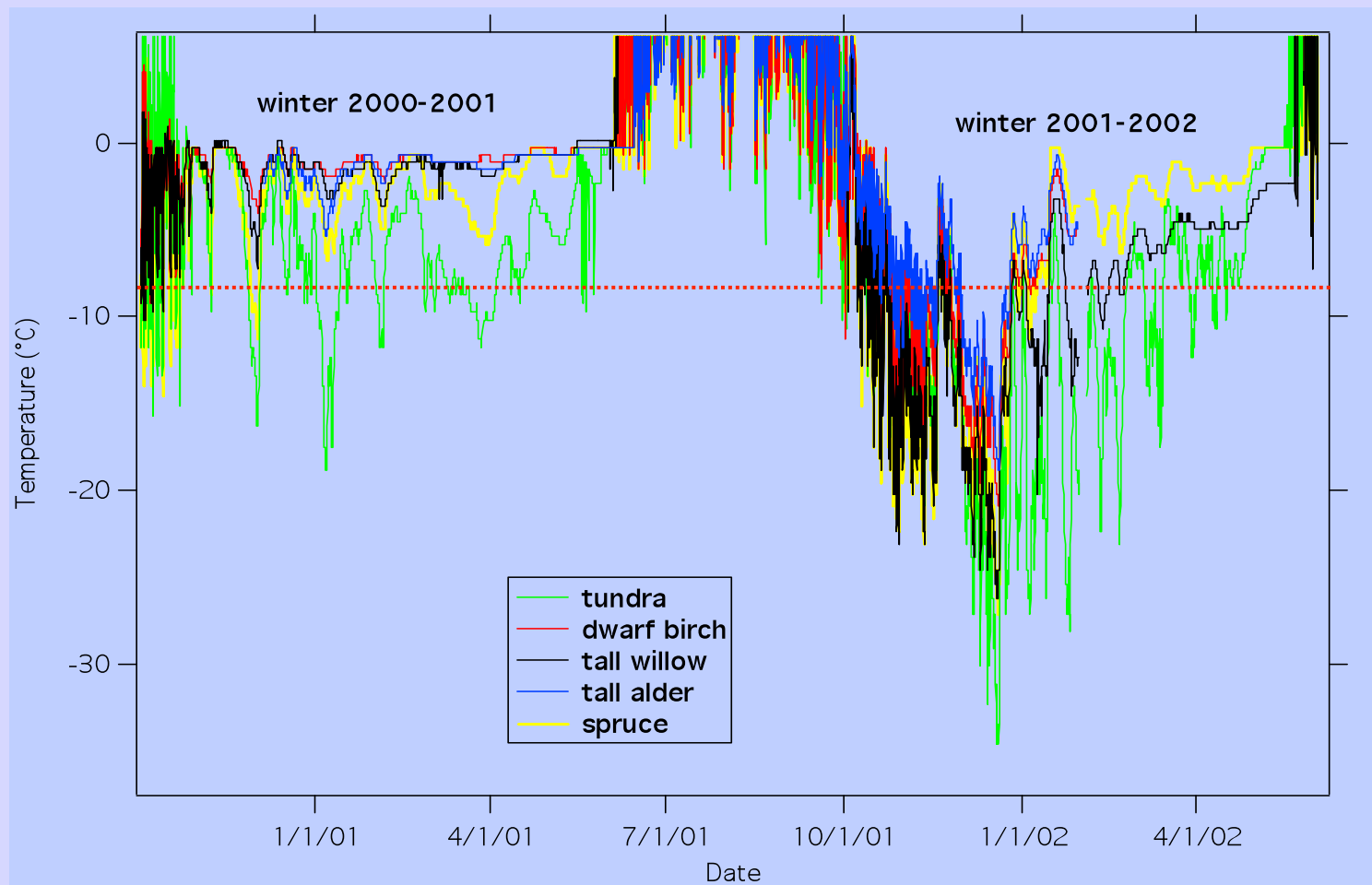
The way winter comes to the Arctic-permanent snow pack arrival dates vary and can have important consequences for ecosystem function

# Seward Peninsula

## 2000-2002: Two very different snow histories



## Variation in ground temperatures due to late, shallow snowpack in 2001-2002.



# Herbivory in winter



# Timing of snow pack development and browsing interactions

## Early snow pack

## Late snow pack

### Abiotic

### Herbivores

Warmer soils



Higher winter N and C mineralization-greater shrub growth in summer



Larger shrub canopies, reduced graminoids, reduced soil C storage-



Increases in winter and summer albedo?  
Feedbacks?



Reduced winter forage availability



Increases cratering, more energy expenditure, poorer forage quality



Nutritional deficits, reductions in fecundity



**Subsistence harvesting reduced**

### Abiotic

### Herbivores

Colder soils



Lower winter C & N mineralization-reduced shrub growth



Smaller shrub canopies, increased graminoid and lichen growth, increased soil C storage



Decreases in winter & summer albedo?  
Feedbacks?



Greater winter forage availability



Reduced cratering, reduced energy expenditure-



Higher forage quality, increases in fecundity

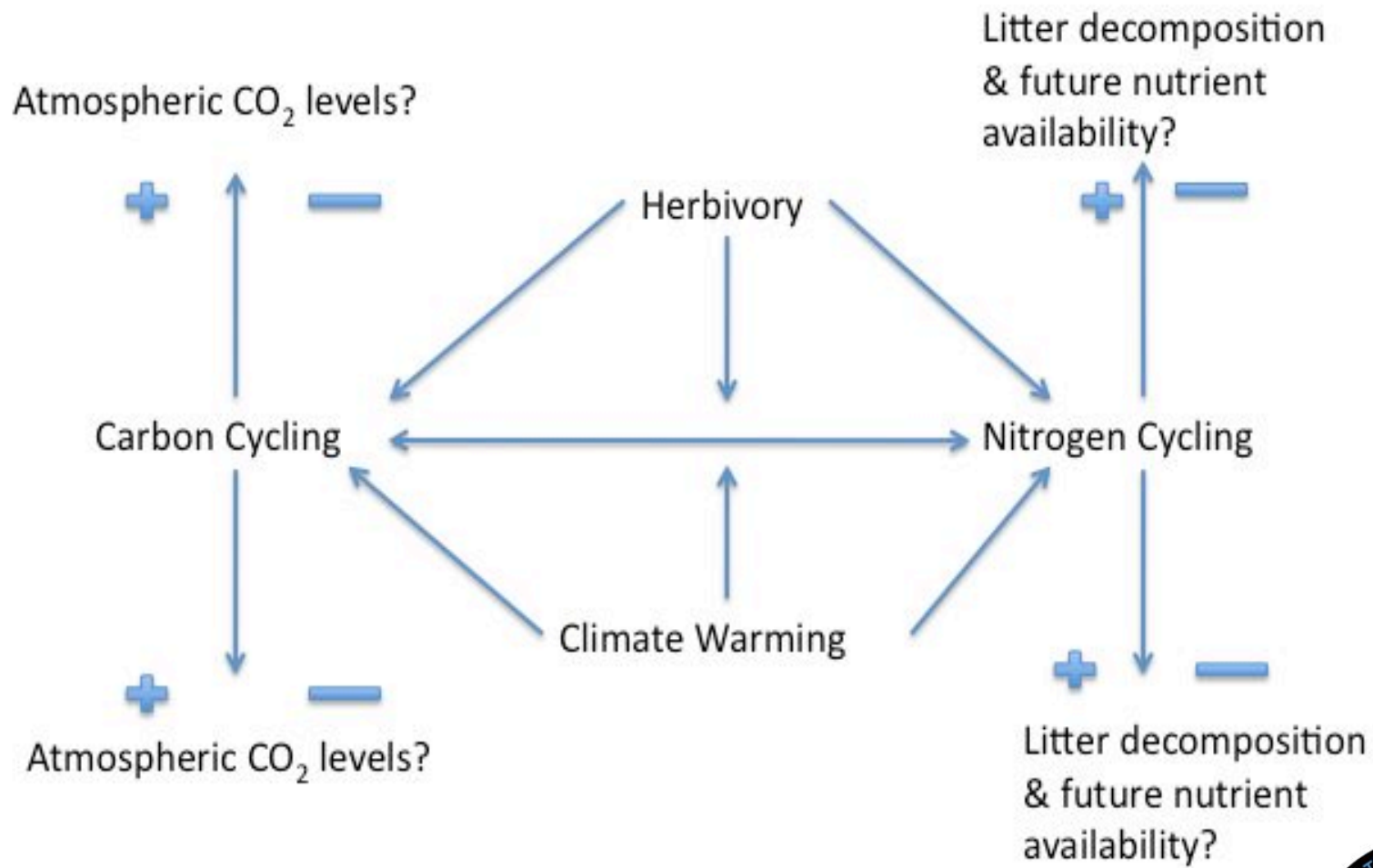


**Subsistence harvesting increased**





Atmospheric CO<sub>2</sub> levels?



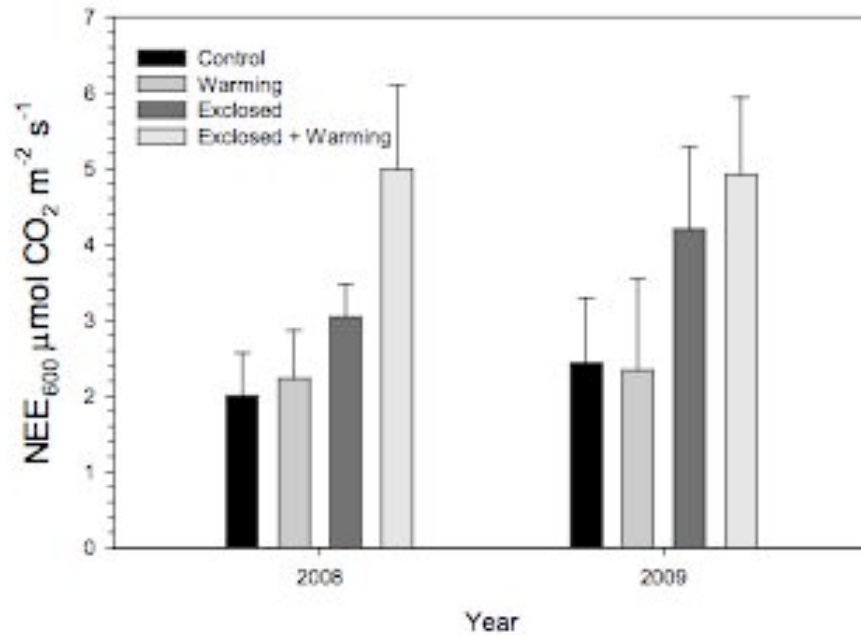


Measurements:  
CO<sub>2</sub> flux-NEE, GEP, RE  
Vegetation canopy traits  
Plant leaf C&N,  
Soil C & N  
Vegetation phenology  
Soil Moisture

Treatments: Exclosed-ambient  
and warmed, Grazed-ambient  
and warmed: started in 2003



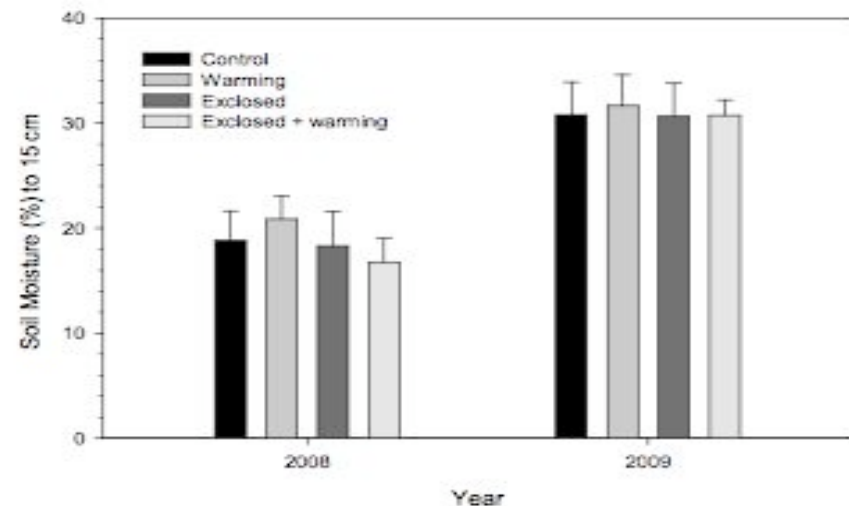
NEE (normalized to  $600 \mu\text{mol photons m}^{-2} \text{s}^{-1}$ ) at peak season for both years



Warming response is muted when herbivores are present and are stronger in dry vs. wet years

The uncertainty is however, the extent to which grazers effect belowground processes, and the effects on the synchrony of growth between roots and shoots

Interannual comparison of soil moisture at peak season (mid-July)



**Soil C in the High Arctic**



## Soil Carbon Pool Reassessment of the High Arctic

Vegetation type	Bliss and Matveyeva (1992)			Horwath and Sletten (this work)	
	Area (10 <sup>6</sup> km <sup>2</sup> )	O.M. (kg/m <sup>2</sup> )	SOC (Pg)	Organic Carbon (kg/m <sup>2</sup> )	SOC (Pg)
Polar semidesert	1.005	2.19	0.978	8.72	8.68 <sup>a</sup>
Polar desert	0.847	0.046	0.017	2.52	2.13 <sup>a</sup>
Mire	0.132	23.36	1.358	29.8	3.93 <sup>a</sup>
Total			2.35		14.74

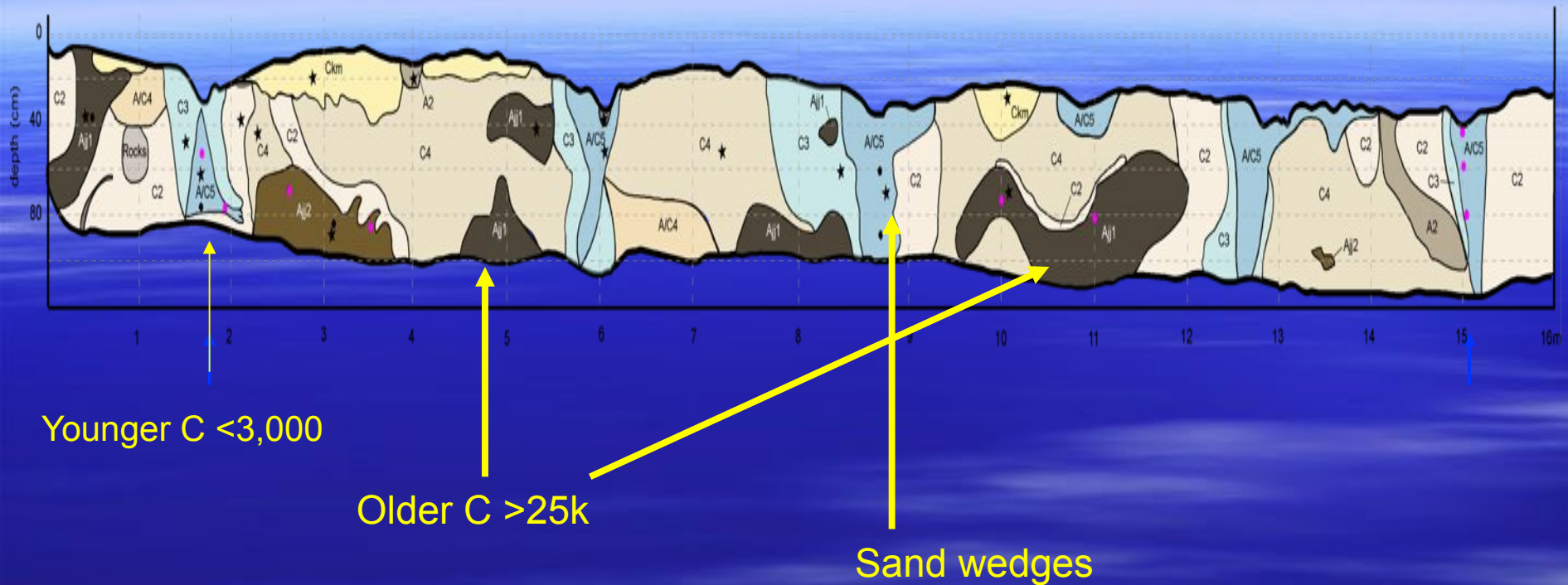
~ 8 X more carbon estimated for High Arctic Systems



# Trenching of polar stripe area for detailed soil carbon studies



# Polar strip formation, soil carbon distribution and ages ( $^{14}\text{C}$ dating)



- Vegetated troughs are filled with sand, indicating importance of aeolian processes formation and stabilization
- Soil C ages increase with depth in sand wedges
- Older soil C found in pockets that appear to have undergone cryoturbation
- Two distinct pools of carbon – younger “active” and older “preserved”



A wide-angle photograph of a polar landscape. The foreground is a vast, flat expanse of brownish-grey soil and rocks, with scattered patches of snow. The middle ground shows a similar terrain with more snow patches. In the background, there are low, rolling hills and mountains, some with snow on their peaks. The sky is overcast and grey.

## Key Questions

What are the magnitudes of soil respiration over the summer in polar stripe landscapes?

Is there evidence that the older soil carbon is contributing to the total soil C efflux?

# Measuring and Sampling of CO<sub>2</sub>

Rate of CO<sub>2</sub> flux from the Soil Surface (Soil Respiration)



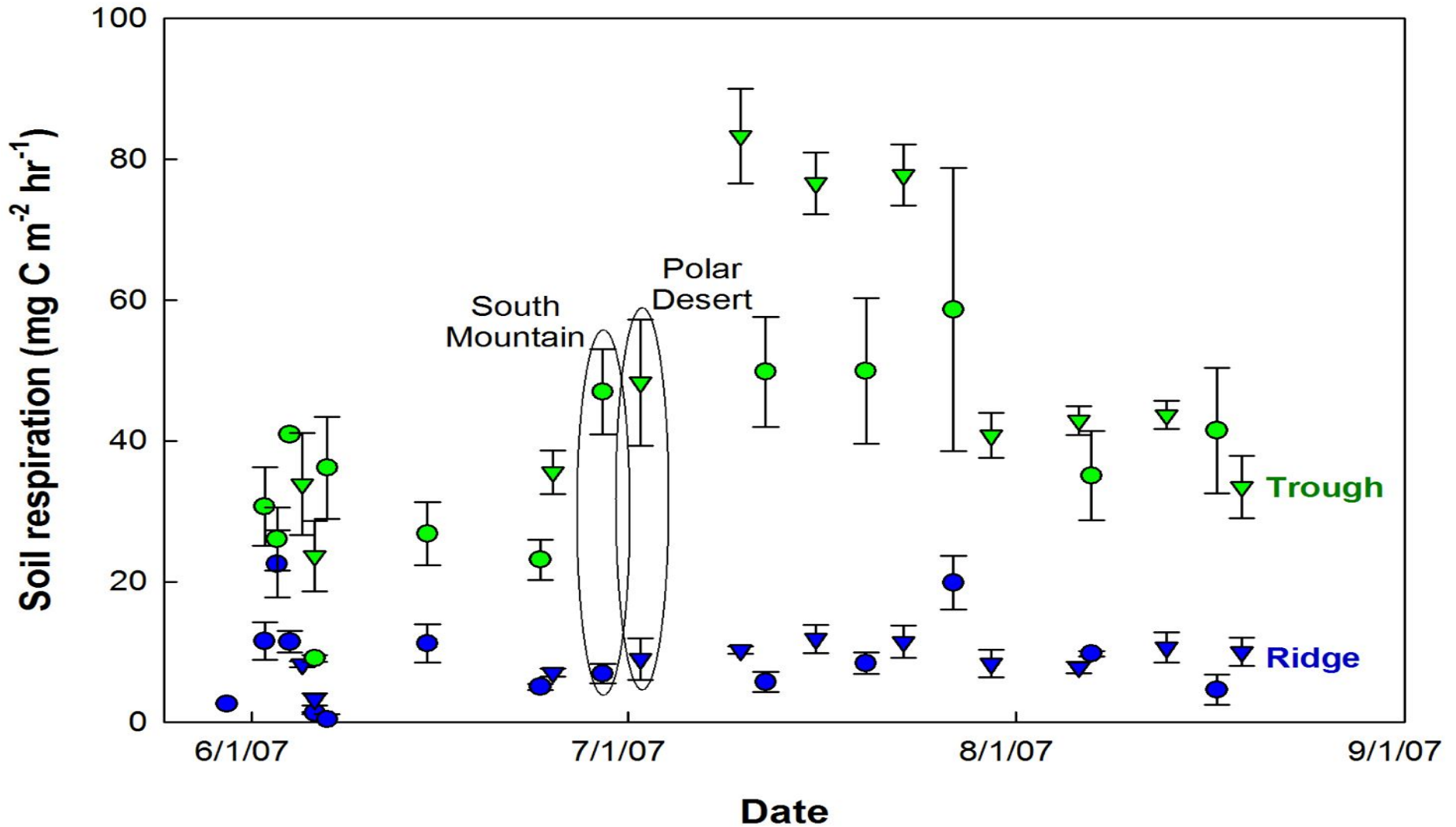
Storage of  
air samples

Soil  
Probes



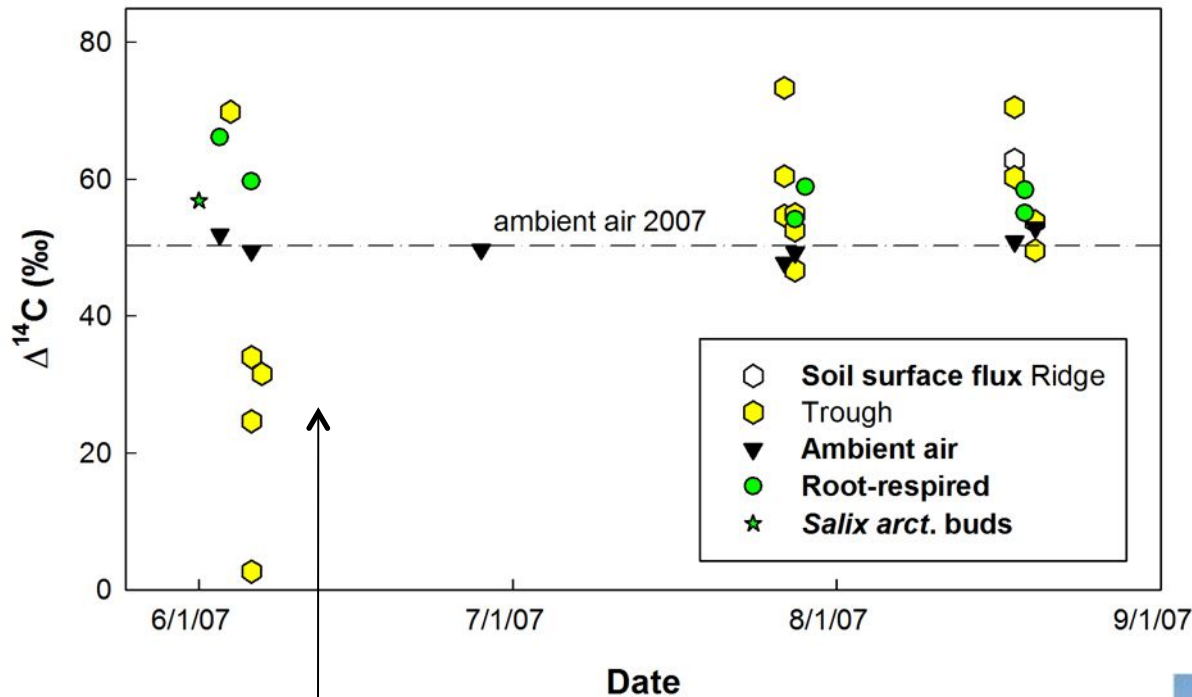
[CO<sub>2</sub>]- Depth Profiles





Soil CO<sub>2</sub> efflux rates (respiration) are 8x higher on vegetated troughs than barren ridges and show a strong seasonal pattern

# Delta $^{14}\text{C}$ of respired $\text{CO}_2$ from polar stripe landscapes



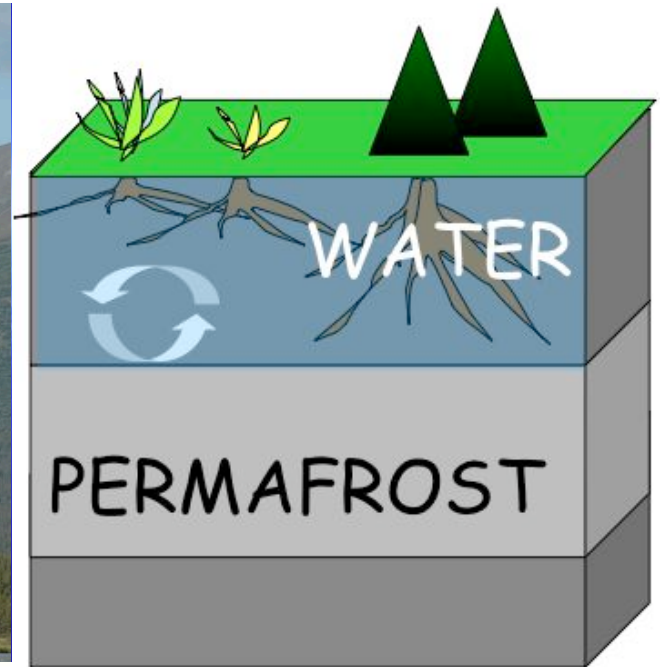
Efflux of older soil C is detectable before leaf out when the signal becomes swamped by fresh C from recent photosynthesis and root exudation

Leaf out

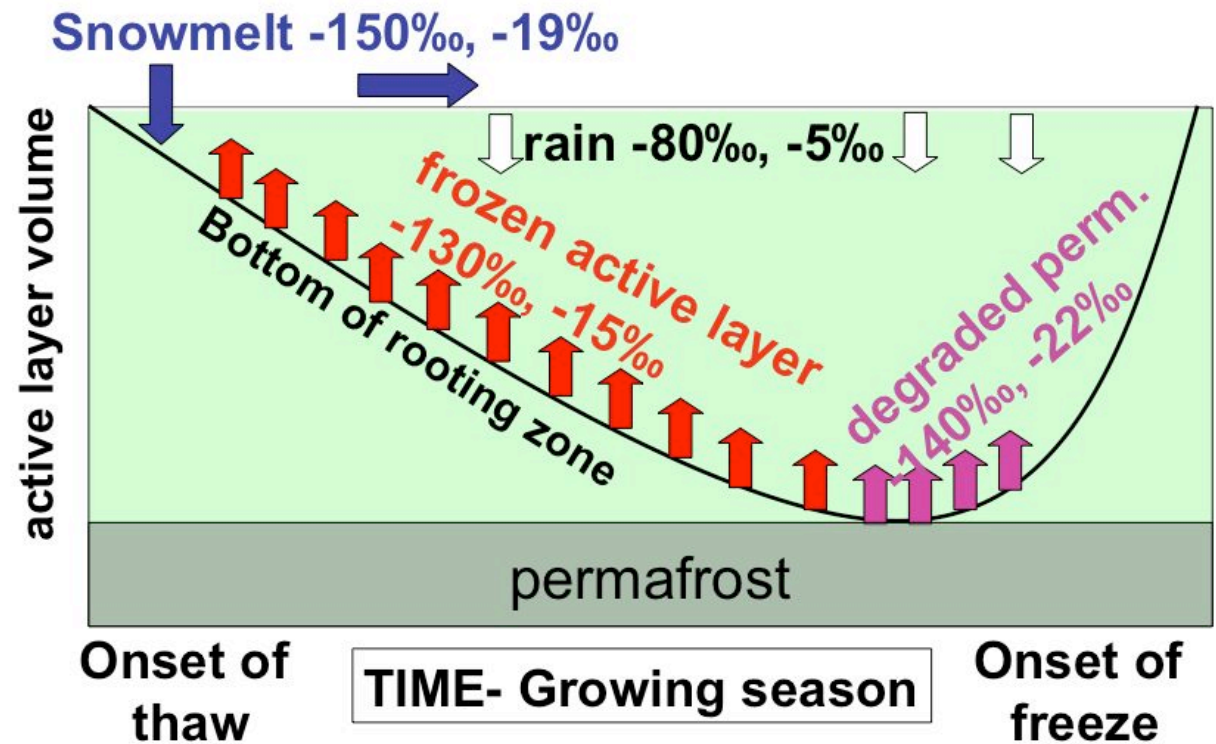
How will older soil C efflux respond to climate warming with and without added summer water?

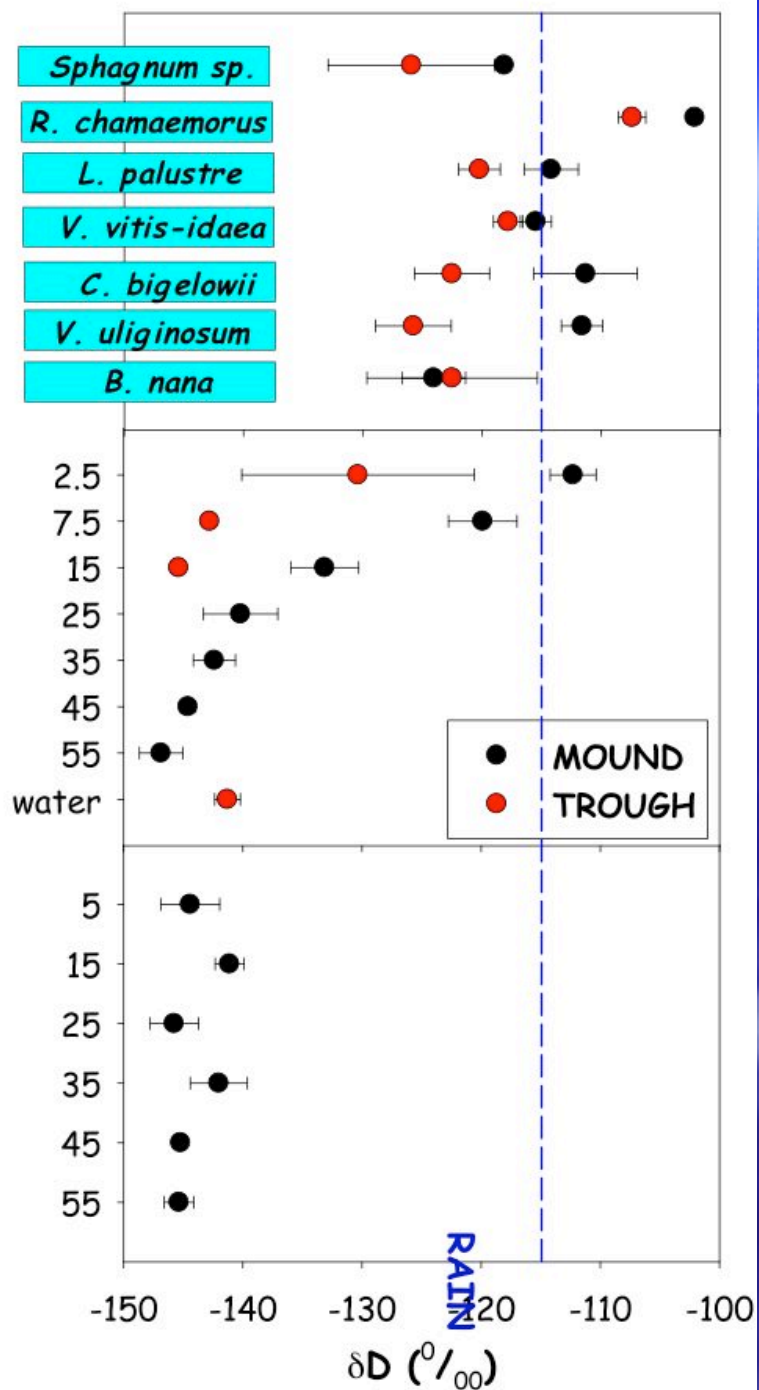


# Permafrost and Ecohydrology



Thawing permafrost may become an important water source for boreal and arctic plants-leading to: vegetation change and changes in the patterns and magnitudes of transpiration





### evergreen



*Ledum palustre decumbens*



*Picea mariana*



*Vaccinium vitis-idaea*

### deciduous



*Betula nana*



*Vaccinium uliginosum*



*Salix arbusculoides*



*Salix pedicellaris*

Plant use of rain and deep water vary between life forms, the  $\delta D$  values of water in the active layer becomes progressively depleted from the surface to the bottom of the active layer/top of the permafrost.

Uncertainty-how do water sources vary between wet and dry years and as permafrost thaws, will "old" water become an important water source for all species?

## Summary

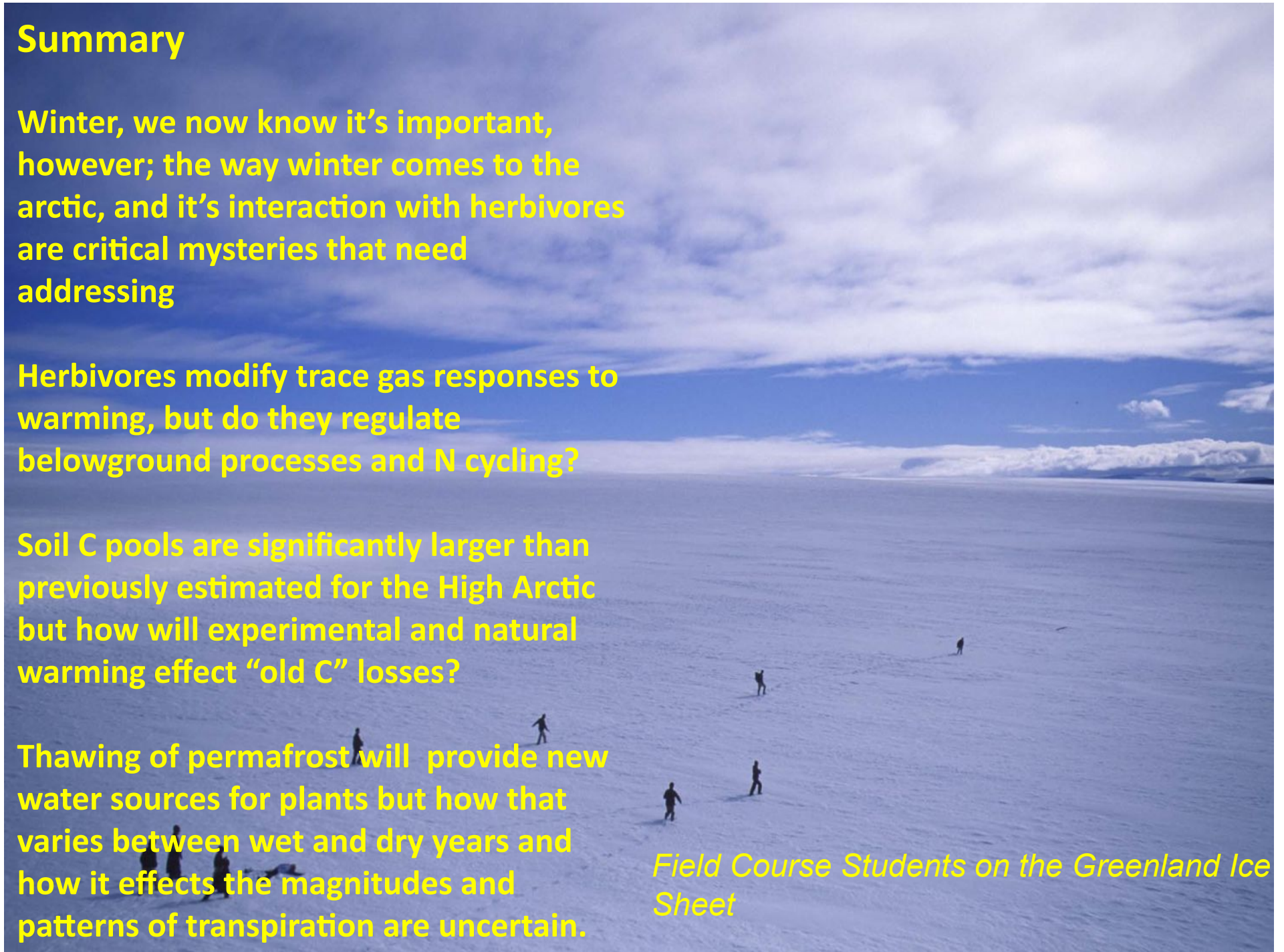
Winter, we now know it's important, however; the way winter comes to the arctic, and it's interaction with herbivores are critical mysteries that need addressing

Herbivores modify trace gas responses to warming, but do they regulate belowground processes and N cycling?

Soil C pools are significantly larger than previously estimated for the High Arctic but how will experimental and natural warming effect "old C" losses?

Thawing of permafrost will provide new water sources for plants but how that varies between wet and dry years and how it effects the magnitudes and patterns of transpiration are uncertain.

*Field Course Students on the Greenland Ice Sheet*



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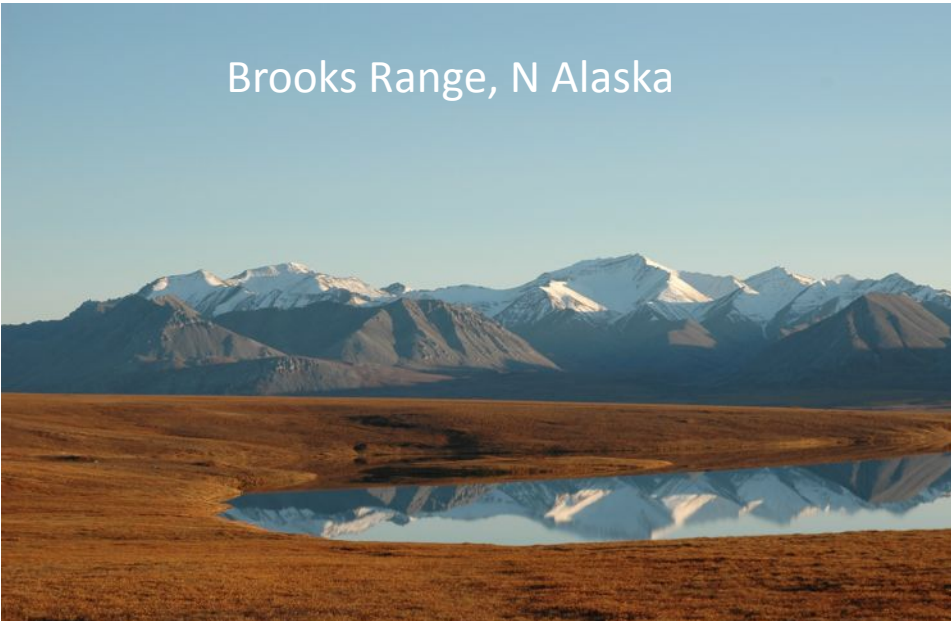
Alyeska in March



Ripping the Aly pow



Brooks Range, N Alaska



West Central Greenland



NW Greenland

